Calculation of Diffracted Multiple Reflections

by Jon F. Claerbout and Don C. Riley

Prerequisite reading material is "Two Techniques for Wave Equation Migration" also contained in this report.

For downward continuing a downgoing wave we have the equation

$$D_{zt}, = \frac{c}{2} D_{xx}$$
 (1)

which is equations (8) or (15) in the Claerbout & Johnson paper. In equation (1) the independent variable x is a receiver location. Equation (1) can be converted to a form for use on zero offset sections if we define the shot-receiver midpoint y as y = x/2. This gives

$$D_{zt}, = \frac{c}{8} D_{yy}$$
 (2)

Let T be the tri-diagonal matrix with - c  $\Delta z \Delta t$  / 32  $\Delta y^2$  times the second difference operator  $\delta_{yy}$  on its main diagonal. Then equation (2) with the boundary condition that D vanishes before t' = 0 can be written in the tabular form

Given that the initial conditions are that the downgoing waveform is observed at the surface z=0 for times t'=0, 1, 2, and 3 (but not t'=4 and beyond) we can fill in table (3) in the order indicated below ( g denotes something given as an initial condition)

		z			
	0	0	0	0	0
	g	1	3	6	10
	g	2	5	9	
	g	4	8		
t'	g	7			
¥	?				

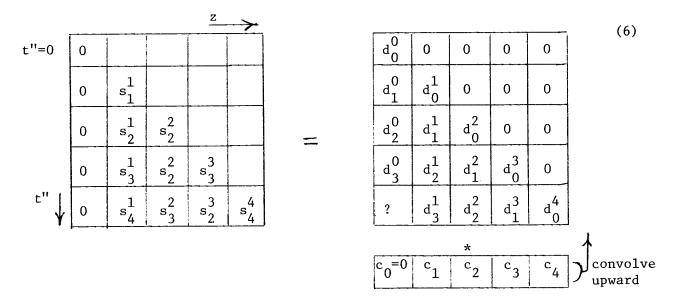
Down-shifting successive columns in the table (4) we are converting t' to a t" coordinate.

	z=0			Z	<b>&gt;</b>
	0	0	0		
t"=0	$d_0^0$	0	0	0	0
	d <sub>1</sub> 0	$d_0^1$	0	0	0
	d <sub>2</sub> <sup>0</sup>	$d_1^1$	$d_0^2$	0	0
. 11	d <sub>3</sub> <sup>0</sup>	$d_2^1$	d <sub>1</sub> <sup>2</sup>	d <sub>0</sub> <sup>3</sup>	0
t"	?	$d_3^1$	$d_2^2$	d <sub>1</sub> 3	d <sub>0</sub> <sup>4</sup>

The meaning of the t'' coordinate is that all elements on a given row of constant t'' can contribute to the upcoming received wave at a surface arrival time t''. Now equation (11) in the section "Programming Techniques for Wave Equation Migration" modified by y = x/2 reads

$$\partial_{tz}$$
 U(y,z,t) =  $-\frac{\overline{c}}{8}$   $\partial_{yy}$  U(y,z,t) - c(y,z)  $\partial_{t}$  D(y,z,t-2z/ $\overline{c}$ )

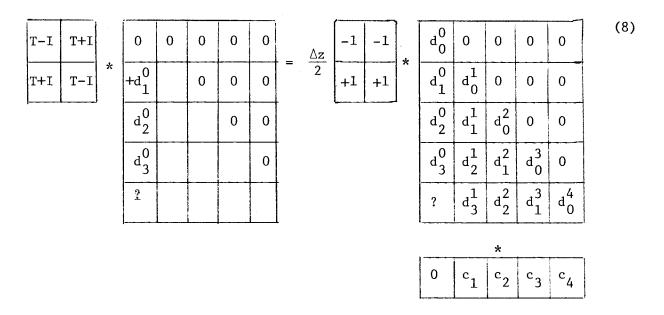
The right hand term which is the source term S(y,z,t) acting in a region  $\Delta z$  for up-coming waves can be expressed in table form as



In (6) we have set the surface reflection coefficient  $c_0 = 0$  because we will handle the surface boundary condition D = -U separately. This means that "?" drops out of the left side of (6). The negative of equation (5) is

$$(\frac{\overline{c}}{8} \partial_{yy} + \partial_{tz}) (-U) = c \partial_{t} D$$
 (7)

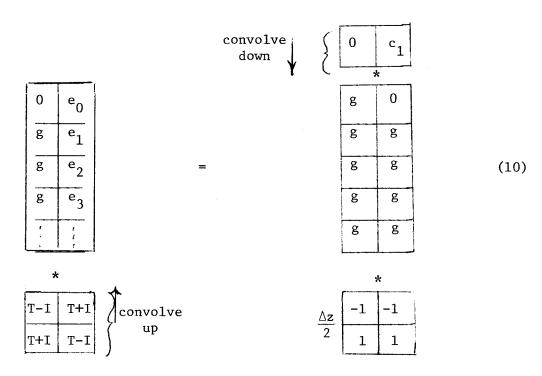
which can be written in table form as



Now suppose the  $-\mathtt{U}$  table was completely specified for all z and all those t" which are at and before t" = 3. Now we can calculate '?' =  $\mathrm{d}_4^0$  from the bottom two rows of (8) by calculating the entire bottom row in the order indicated

All the steps up to the present have shown how knowledge of all the reflection coefficients and the surface wave D = -U for times t = 0, 1,2,3 can be used to calculate the surface waves at time t = 4, namely  $d_4^0 = -u_4^0$ . Obviously the process may be used recursively to get  $d_t^0 = -u_t^0$  for all time.

The inverse calculation proceeds in a similar fashion. Suppose  $d_t^0$  is known for all t and we wish to calculate  $c_1$ ,  $c_2$ , etc., in a recursive fashion. It is sufficient to show how to compute  $c_1$ . Suppose we skim off the left two columns of (8). We get



In (10) all the boxes filled by g on the left are given. The boxes with g on the right are readily computable as before. If  $c_1$  were known it would be a straightforward task to compute successively ...  $e_3$ ,  $e_2$ ,  $e_1$ ,  $e_0$ . We would be compelled to initialize the computation with an approximation such as  $e_N = 0$  for some large N . If the correct value of  $c_1$  had been used then we should find that  $e_0$  vanishes. Since we do not know what value of  $c_1$  to use we try  $c_1 = +1$  obtaining  $e_0^+$  and we try  $c_1 = -1$  obtaining  $e_0^-$ . The correct value of  $c_1$  is the appropriately weighted linear combination

$$0 = \alpha e_0^+ + \beta e_0^- \tag{11}$$

where

$$1 = \alpha + \beta \tag{12}$$

$$c_1 = \alpha - \beta \tag{13}$$

which inverts to

$$2 \alpha = 1 + c_1 \tag{14}$$

$$2\beta = 1 - c_1 \tag{15}$$

reducing (11) to

$$0 = (1 + c_1) e_0^+ + (1 - c_1) e_0^-$$

or

$$c_1 = \frac{e_0^+ + e_0^-}{e_0^+ - e_0^-} \tag{16}$$