

INVERSION OF REFRACTION DATA BY WAVE-FIELD MIGRATION

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Abstract

The process of wave equation migration is adapted for refraction data in order to produce velocity-depth models directly from the recorded data. The procedure consists of two linear transformations: a slant stack of the data produces a wave field in the $p-r$ plane which is then migrated using $r=0$ as the imaging condition. The result is that the data wave field is linearly transformed from the time-distance domain into the slowness-depth domain, where the velocity profile can be picked directly. No traveltimes picking is involved, and all the data are present throughout the inversion.

The method is iterative because it is necessary to specify a velocity function for the migration. The solution produced by a given iteration is used as the migration velocity function for the next step. Convergence is determined when the migrated wave field images the same velocity-depth function as was input to the migration.

The method obviates the problems associated with determining the envelope of solutions that are consistent with the observations, since the time resolution in the data becomes transformed into a depth resolution in the slowness-depth domain.

Introduction

Refraction profiles are conventionally analysed by extracting traveltime information from the data and performing a Wiechert-Herglotz integration to produce a velocity-depth profile. In this paper we present an alternate approach to refraction inversion which consists of transforming the entire data wave field into the slowness-depth domain. The process involves two linear transformations: a slant stack followed by a depth migration.

Examples of the processing of refraction wave fields by slant stacking were presented by McMechan and Ottolini (1980), Stoffa and Buhl (1979) and Phinney et al. (1980). The result of this transformation is a wave field in the ray parameter-time intercept (p - τ) plane. The image that forms in the p - τ plane is the "tau" curve (cf. Bessonova et al. 1976), required for Wiechert-Herglotz inversion. The reverse of this process was previously used by Chapman (1978) to generate synthetic seismograms from a p - τ curve.

Recently, Garmany et al. (1979) showed that the inversion of a tau curve can be expressed in linear form if the integration is performed along lines of constant p . In this paper we employ a migration method to linearly transform from the p - τ domain directly into the slowness-depth domain. Since both slant stacking and migration transform the entire wave field, neither the traveltime curve nor the tau curve needs to be picked. Instead, picking is delayed until the last step when the slowness-depth model is extracted directly from the migrated wave field. The wave-field transformation approach to inversion has the advantage that all the data contribute to the final image; there is no subjective selection of data (e.g. via traveltime picking) as is involved in previous methods. In theory, since both migration and slant stacking are reversible transformations, this approach could potentially be used to generate synthetic refraction gathers from the p - z plane.

The migration part of this technique is iterative because it is necessary to specify a velocity function. Convergence is achieved when the extracted velocity function is the same as that input to the migration. In the examples we have tested, convergence was obtained in less than five iterations.

In this paper we present the theory for the inversion of refraction profiles by double transformation of the data wave field. The method is

illustrated with several synthetic examples and with a line of refraction data recorded in the Imperial Valley, California.

Theory

The data recorded in a typical refraction experiment is in the format of a common shot gather. Making the assumption that velocity varies only with depth, it is equivalent to treat such data as a common mid-point gather. We shall use the symbol t for traveltime and h for half offset in a common mid-point coordinate system. The inversion of an observed refraction wave field in these coordinates involves two linear transformations: a slant stack and a depth migration. Each of these will now be considered in turn.

A slant stack is a linear operation that transforms a wave field in the t - h domain into a wave field in the p - τ domain. Here, p is the ray parameter (horizontal slowness), and τ is the vertical component of t [t projected to zero offset along a line of slope p through the point $(t, 2h)$]. Slant stacking of refraction profiles has been discussed in detail by McMechan and Ottolini (1980) so only a brief summary is given here.

Slant stacking can be performed in either the time or frequency domain; however, in the time domain, variable trace spacing is easier to handle. A slant-stacked wave field is produced from a common shot gather by (Gel'fand et al., 1966; Chapman, 1978; McMechan and Ottolini, 1980):

$$S(\tau, p) = \int_{-\infty}^{+\infty} P(\tau + 2ph, h) dh \quad (1)$$

where P is the observed (seismogram) wave field and S is the transformed (p - τ) wave field. Equation (1) can be cast in the frequency domain by using the Fourier central slice theorem

$$S(\omega, p) = P(\omega, -2\omega p) \quad (2)$$

In other words, the two-dimensional Fourier transform of P evaluated along the slice $-2\omega p$ is the Fourier transform with respect to time of its projection

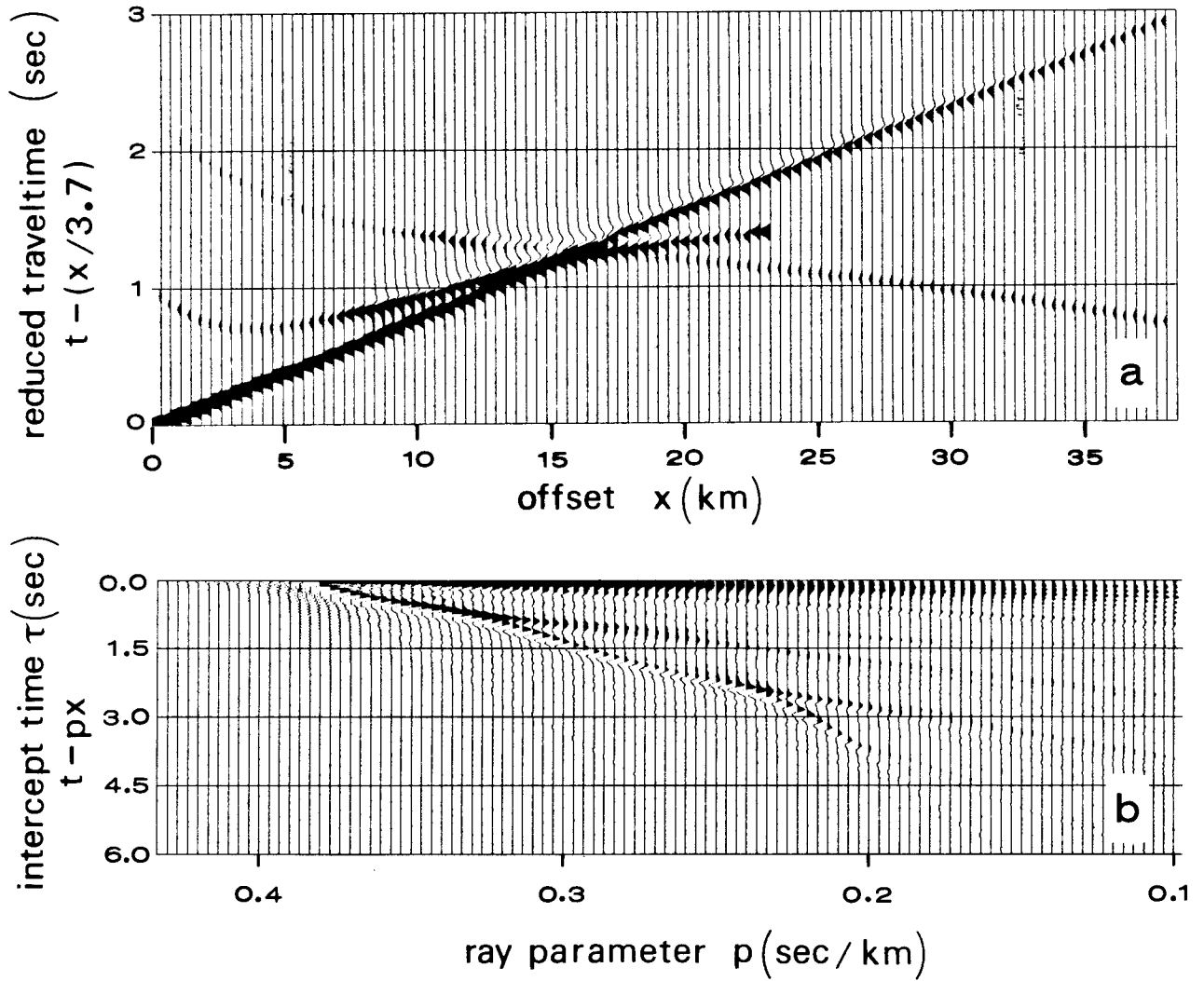


FIG. 1. The first transformation. The upper (a) half of this figure contains a synthetic refraction profile plotted in reduced time format. The reduction velocity is 3.7 km/s. These data are transformed by slant stacking into the plane wave decomposition shown in the lower (b) half of the figure. This transformation is the first half of the process of inversion of the data wavefield. The result of migrating this slant stacked wavefield (b) is shown in Figure 2.

$S(\tau, p)$. For a more detailed presentation of these concepts see Bracewell (1956) and Bracewell and Riddle (1967).

The main result of slant stacking is that the observed wave field is decomposed into plane wave (fixed p) components, each of which can be migrated separately. Figure 1a contains an example of a synthetic common shot gather, and 1b is the corresponding slant stack. The finite aperture of the dataset and spatial aliasing both contribute to artifacts in the slant stack. Relatively dense sampling of the offset coordinate is required to produce a reasonable image in the p - τ domain (cf. McMechan and Ottolini, 1980). Also, coherency of source signatures is required if multiple shots are involved in the field survey.

The second transformation in this technique is a depth migration of the slant-stacked wave field. Migration can be considered in two parts: a downward continuation of the wave field followed by the application of an imaging principle. When velocity varies only with depth [$v = v(z)$], the downward continuation of the wave field observed at the surface ($z = 0$) to any desired depth (z) can be implemented by a phase rotation in the frequency domain (cf. Claerbout, 1976; Gazdag, 1978):

$$P(\omega, k_h, z) = P(\omega, k_h, 0) \exp \left[-i2 \int_0^z \left(\frac{\omega^2}{v^2(z)} - \frac{k_h^2}{4} \right)^{1/2} dz \right] \quad (3)$$

where ω is the temporal frequency and k_h is the horizontal wavenumber (the dual of h). Equation (3) is an approximate solution to the wave equation

$$\left(\frac{\partial^2}{\partial z^2} + \frac{\partial^2}{\partial h^2} + 4 \frac{\omega^2}{v^2(z)} \right) P(\omega, h, z) = 0 \quad (4)$$

The solution is approximate because amplitude terms depending on the logarithmic gradient of the velocity have been neglected. However, the traveltimes aspects of the solution are correct. The first minus sign in equation (3) indicates that we are imaging upcoming waves.

To convert equation (3) into slowness form, substitute $-2\omega p$ for k_h :

$$P(\omega, -2\omega p, z) = P(\omega, -2\omega p, 0) e^{-i\Psi(\omega, p, z)} \quad (5)$$

where

$$\Psi(\omega, p, z) = 2\omega \int_0^z \left[v^{-2}(z) - p^2 \right]^{1/2} dz \quad (6)$$

Using relation (2), equation (5) may be rewritten in terms of slant-stacked wave fields:

$$S(\omega, p, z) = S(\omega, p, 0) e^{-i\Psi(\omega, p, z)} \quad (7)$$

The inverse Fourier transform of equation (7) is

$$S(\tau, p, z) = \int S(\omega, p, 0) e^{-i\Psi(\omega, p, z)} e^{-i\omega\tau} d\omega \quad (8)$$

With equation (8), the surface ($z=0$) slant stack can be extrapolated (downward continued) to all depths. One could think of using this equation to fill out the entire τ - p - z space with data extrapolated from the surface. However, this is not necessary because the slowness image we seek lies on a plane in the τ - p - z space specified by the imaging condition. The desired image is that which contains the bottoming points of all the plane wave components. The plane which images this trajectory is specified by the condition $\tau = 0$ because we wish to downward continue each p to the depth at which the corresponding ray bottoms. At that depth, p of the ray equals the true slowness (v^{-1}) of the medium for refractions and postcritical reflections. Pre-critical reflections image in a trajectory that splits off from the main slowness image at the critical reflection point. Setting $\tau = 0$ in equation (8) yields the desired result:

$$S(0, p, z) = \int S(\omega, p, 0) e^{-i\Psi(\omega, p, z)} d\omega \quad (9)$$

Implementation of equation (9) is straightforward. Each plane wave component (each p) can be migrated separately since p enters equation (9) as a parameter. A computer program which realizes (9) consists basically of three nested loops, an outer one over p , then one over z , and an inner one over ω . One minor problem that occurs in the application of equation (9) is that Ψ has a branch cut. We remedied this by altering the definition of Ψ to be

$$\Psi(\omega, p, z) = 2\omega \int_0^z \left| v^{-2}(z) - p^2 \right|^{\frac{1}{2}} dz$$

This stops the migration from attenuating the wave field below depths where p is greater than v^{-1} . The application of equation (9) produces an image in a wave field in p - z space. The image trajectory is composed of wavelets whose shapes are each defined by the phase shift associated with the reflection coefficient at each z . A refracted ray can be treated as having an effective reflection coefficient of $-i \operatorname{sgn}(\omega)$ (Chapman, 1978).

Figure 2a shows the p - z wave field obtained by migrating the p - r wave field in figure 1b using the $v(z)$ function (the solid line) superimposed on figure 2a. This $v(z)$ function used for migration is the same as that used to generate the synthetic seismograms (figure 1a). The migrated wave field in figure 2a therefore images the corresponding $v(z)$ function exactly. The varying offset between the input $v(z)$ function and the image is due to the fact that there are phase shifts related to reflection coefficients which are not yet taken into account. Accurate location of the image trajectory is discussed below.

The object of inversion is to find the $v(z)$ [or, equivalently, in the present formulation, the $v^{-1}(z)$] function. However, the migration step requires a velocity-depth function. Thus, obtaining a solution is necessarily an iterative process. Convergence is presumably to a unique model because there are no artificial constraints on the convergence path; the dataset itself contains and provides the solution. The convergence criterion is that the migrated wave field images the same $v^{-1}(z)$ function that was input to the migration, and this state can be detected by comparing the p - z wave fields at successive iterations.

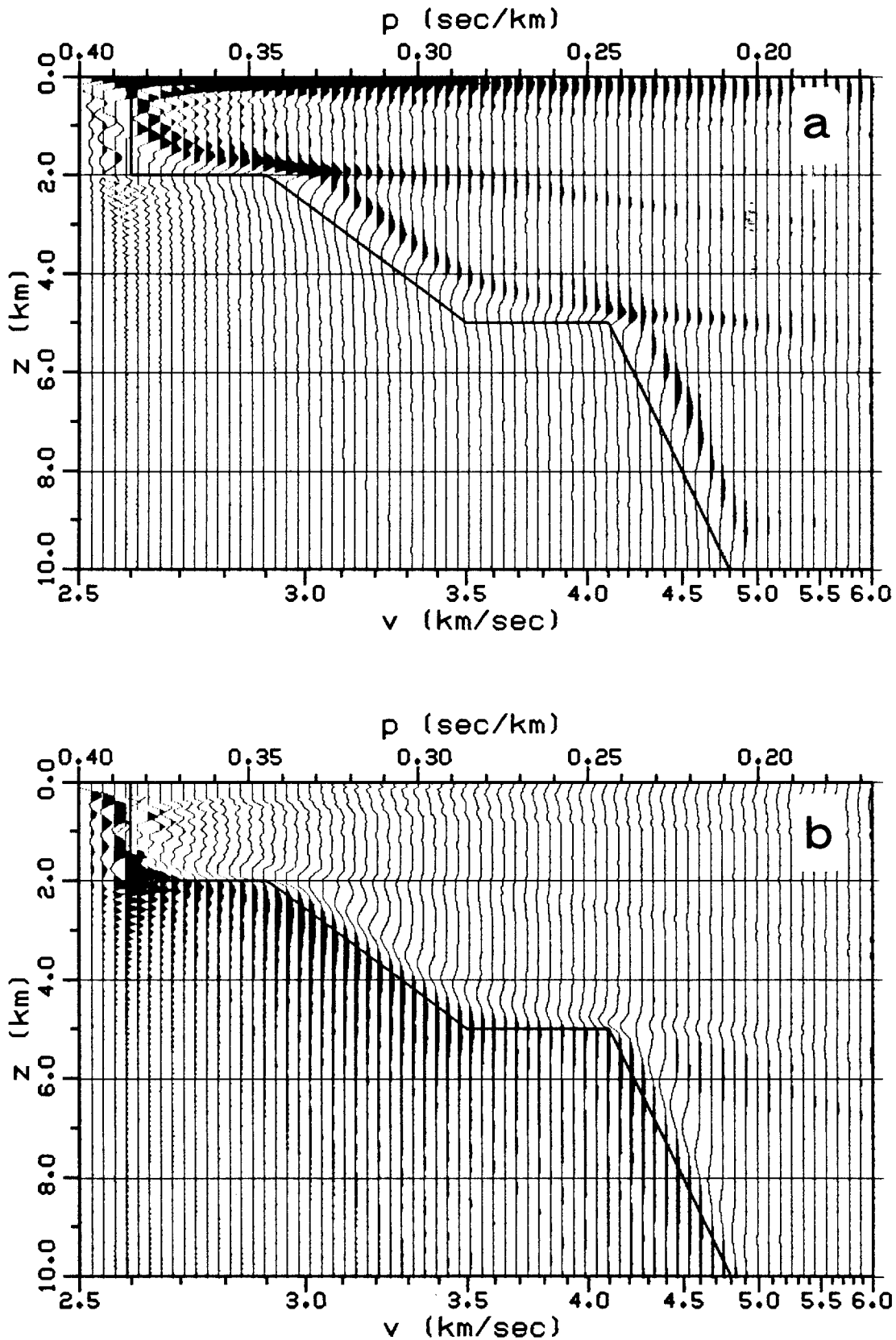


FIG. 2. The second transformation. The upper (a) half of this figure contains the result of migration of the slant stacked wavefield in Figure 1b with the the correct velocity-depth function (the solid line). The lower (b) half of this figure contains the same wavefield as the upper, but with a phase retardation of $5\pi/4$. All the migrated wavefields in the remainder of this paper are presented with this phase shift applied.

Synthetic Examples

It is possible to implement wave field inversion in an interactive mode in which the slowness function extracted from the migrated wave field at each step is used as the migration velocity for the next iteration. In this section we present a detailed example of inversion in an interactive mode. An example of automated convergence is included but the production of a totally automatic inversion scheme that is able to adapt to a variety of conditions is far beyond the scope of this paper.

The most important aspect of inversion by wave-field imaging is the determination of the location of the desired slowness trajectory in the migrated (p - z) wave field. In figure 2a, which illustrates the result of migration with the correct $v(z)$ function, it is clear that the relationship between the image in the wave field and the corresponding slowness trajectory (the solid line) is not trivial. Fortunately, this relationship can be predicted, and as we shall show below, the migration itself has certain self-stabilizing feedback properties which enable a velocity profile to be accurately estimated from the use of any arbitrary criterion for locating the slowness trajectory, provided that it is consistently applied.

In figure 2a it can be seen that the optimal slowness trajectory is near, but not coincident with, the maximum amplitude locus in the migrated wave field. Part of the reason for this is that the p - z wave field images reflection coefficients, and each of these contains a phase shift which depends on the angle of incidence and the velocity gradient at the ray bottoming point. Specifically, if one assigns zero shift to refraction branches and precritical reflections, then wide angle reflections have phase shifts which progress from $\pi/2$ at the critical reflection to π at large offsets, and any ray which touches an internal caustic receives an additional $\pi/2$ shift. The latter occurs, for example, to rays which are refracted in a region of sufficiently high velocity gradient that a triplication is produced in the travelttime curve. The existence of these phase shifts suggests that better convergence could be obtained by identifying those p ranges containing particular types of arrivals and applying appropriate phase shifts. This identification could be made in the p - τ plane since the radii of curvature of refraction and reflection p - τ loci have opposite signs (cf. figure 1b). Alternatively, an

algorithm which is independent of phase shift could be sought.

We refrain from making progressive phase shifts as a function of p in the p - τ wave field, as described above, because this requires one to make a specific interpretation which may bias the migration results. Instead, we have shifted the phase of the entire migrated wave field (after migration) by $5\pi/4$. The result, for the wave field in figure 2a, is shown in figure 2b. This phase shift consists of three contributions. First, there is a $\pi/2$ shift associated with the requirement that the far-field radiation condition be satisfied (cf. Chapman, 1978). Then, there is a $\pi/4$ shift due to the fact that we are dealing with a two-dimensional representation of propagation in three dimensions (Chapman, 1978). These two shifts are exact for all p values. The final shift of $\pi/2$ is the average associated with the range of reflection coefficients expected for the various arrivals (refractions and reflections) in a typical refraction profile.

With the net phase shift of $5\pi/4$ the locus containing the maximum positive amplitudes (or the first significant pulse when considering real data) in a migrated wave field should be everywhere within $\pi/2$ of the correct slowness locus. Using this modified form of the migrated wave fields, it was found to be straightforward to obtain convergence to the neighborhood of the correct model by using a very simple criterion for determining the approximate location of the slowness trajectory at each iteration. The exact phase shift applied is not critical since it does not enter directly into the end product. It simply alters the criterion for locating the optimal slowness trajectory in the p - z wave field.

Migration is a stable operation, even with highly erratic velocity-depth input functions. It does not matter if slowness trajectories are incorrectly located in intermediate iterations; these effects are not cumulative from iteration to iteration because migration at each step is always of the original slant-stacked wave field. Also, it does not matter if the criterion used to find the slowness trajectory changes from one intermediate iteration to the next. In fact we found that, when picking trajectories by hand, the picking criterion evolved from iteration to iteration, and different criteria were used over different p ranges in the attempt to attain convergence. All of this is expected and is a direct consequence of the various phase shifts associated

with different types of arrivals as described above. None of this evolutionary process enters or biases the final solution provided that the identical picking criteria are used on the two successive migrated wave fields used to identify the condition of convergence.

The slant-stacked wave field in figure 1b was chosen to illustrate the convergence of migration. This example was done in an interactive format, in which the slowness trajectory at each iteration was extracted by hand using the criterion of the maximum positive amplitude at each p . The results are shown in figure 3. Beyond iterations 4 and 5 (figures 3c and 3d), the solution did not converge further, but oscillated between the two positions represented by iterations 4 and 5. This oscillation indicates a consistent bias associated with the criterion of maximum positive amplitude used for picking the slowness locus. This oscillation is a useful behavior and will be discussed below. It is encouraging that such a simple criterion produces a result that is close to and exhibits the same general shape as the desired curve. In fact, for many practical purposes, this level of convergence ($< 2\%$ error in velocity when averaged over the profile) is already adequate.

In this section we have demonstrated the wave field inversion of refraction data as implemented in a simple interactive mode, and we have shown that a good estimate of velocity variation with depth is obtained. The solution could presumably be made as accurate as desired since, in the process of model perturbation, the correct phase shifts could begin to be included to refine the differences between successive iterations. Another useful observation is evident in figure 2a. The existence of precritical reflections (the extension of the image through lower slowness values from the critical reflections at the "knees" of the image) indicates the presence of a significant velocity contrast. The use of such additional information would produce better estimates of appropriate slowness trajectories. Additionally, where a true reflection exists, the associated progressive phase shifts are evident in the migrated wave field (cf. figures 2a and 2b). Another subjective aid is inherent in the level of focussing of the slowness image in the migrated wave field. The image becomes better focussed as the correct velocity estimate is approached. The reason that amplitude constraints (e.g. the increased amplitude associated with a critical reflection) enter the solution only in a subjective way is that the present form of wave-field inversion is still based on

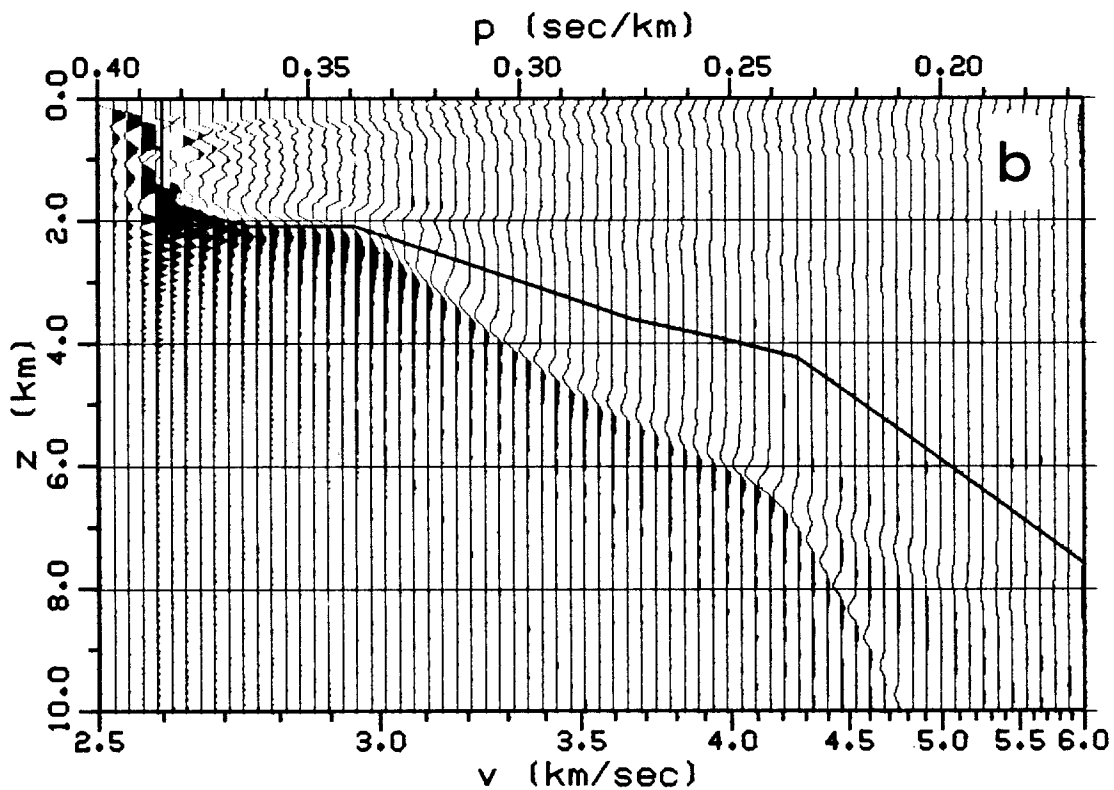
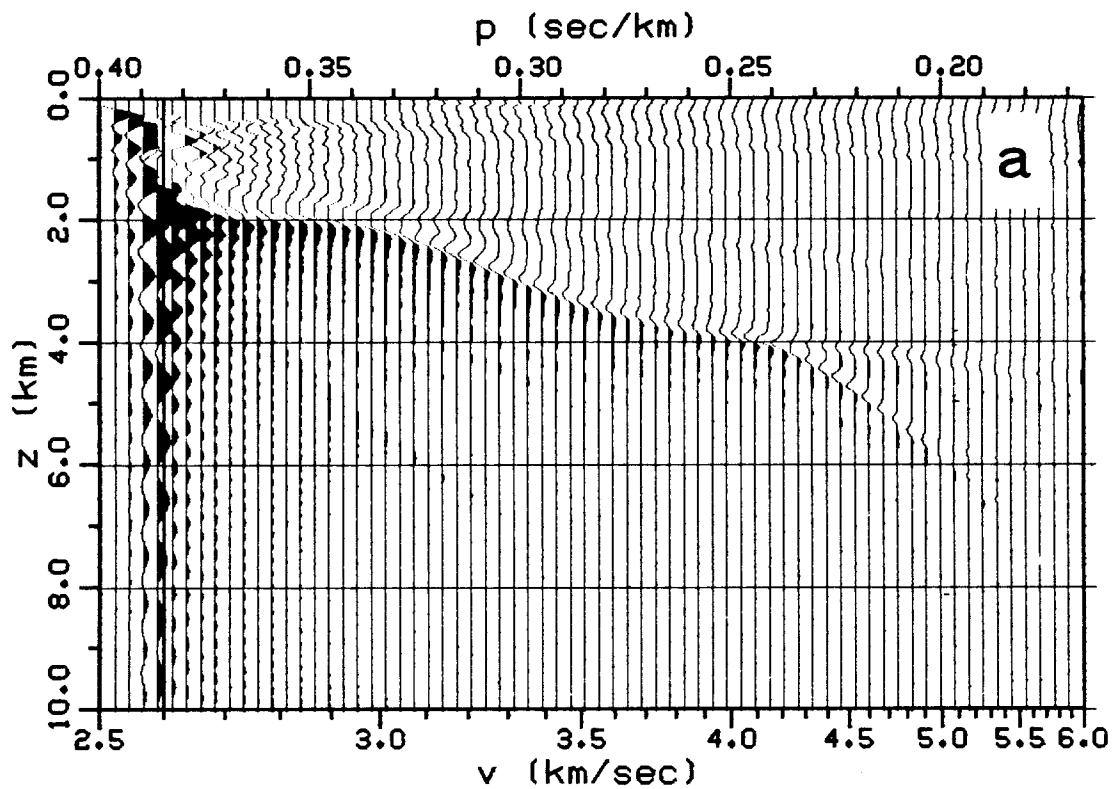


FIG. 3(a,b). Convergence by iteration in an interactive mode. The wavefield being migrated is the plane wave decomposition of the synthetic data shown in Figure 1b. Figure 3a contains the starting $v(z)$ model, a constant velocity of 2.6 km/s at all depths (see the solid line), and the resultant migrated wavefield. The solid line in (b) is the slowness trajectory which was extracted by hand from the wavefield in (a). Migration with this slowness function produced the wavefield in (b). The later stages of this inversion are shown in Figures 3c to 3f.

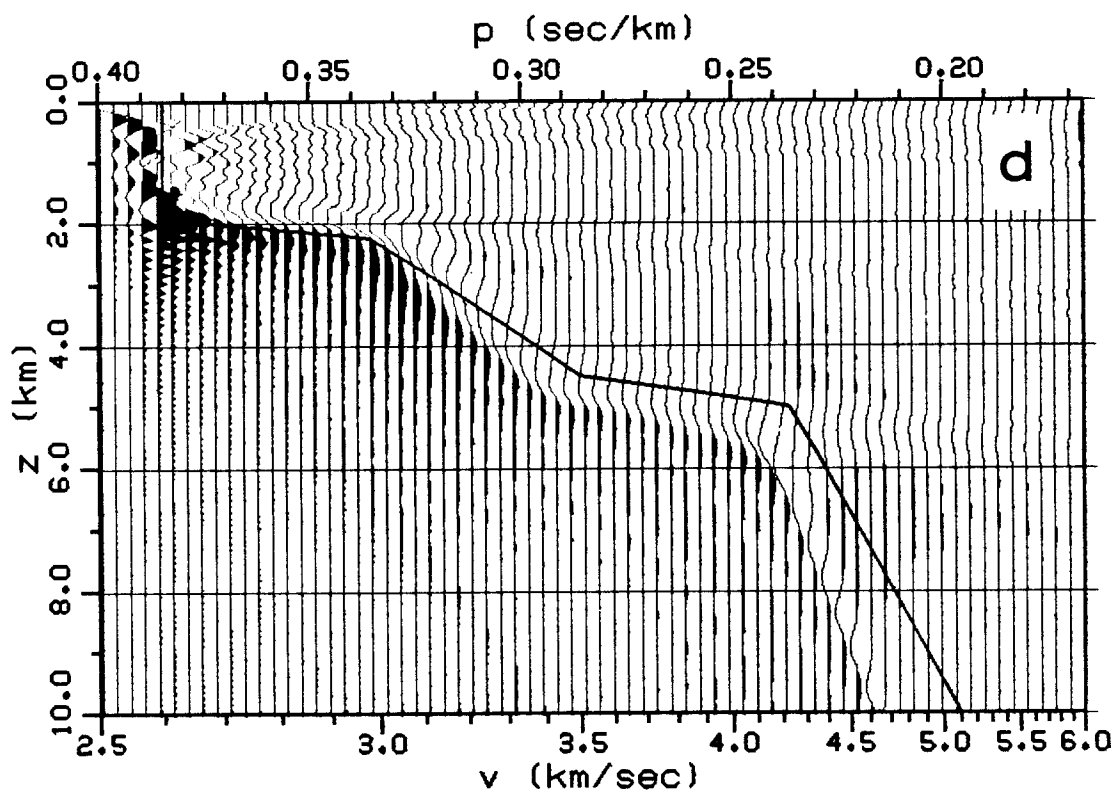
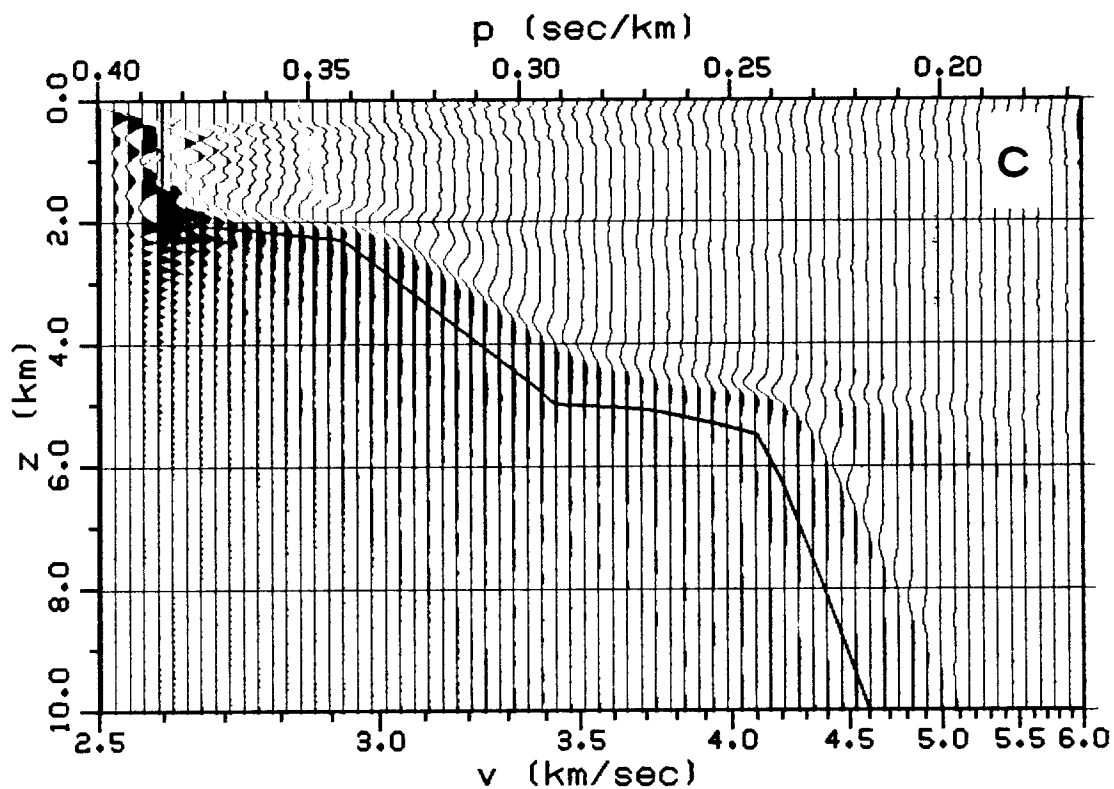


FIG. 3(c,d). This figure is similar to 3a and 3b in that it illustrates intermediate migrations in the process of inversion of the slant stacked wavefield in Figure 1b. Iteration 4 is shown in (c) and iteration 5 is shown in (d). The final results of the inversion are shown in Figures 3e and 3f.

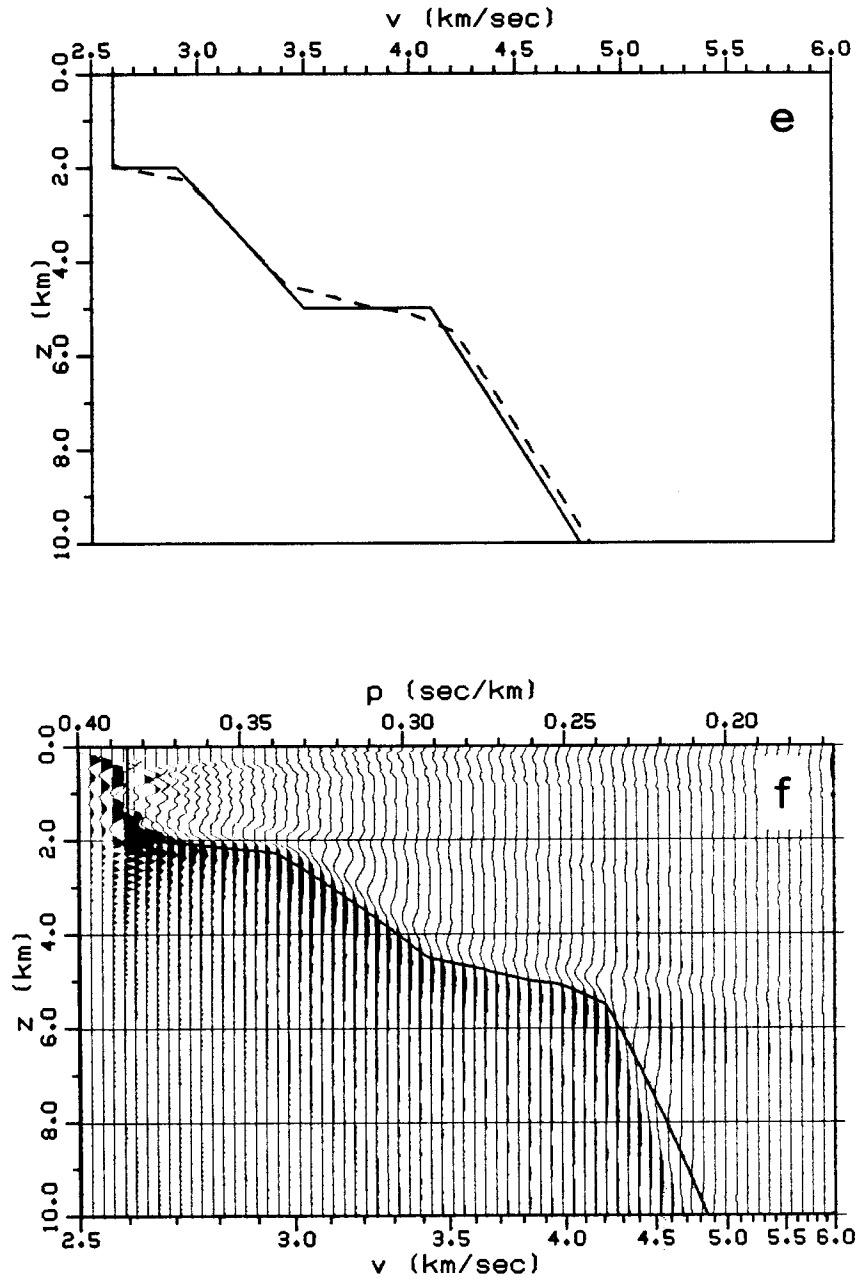


FIG. 3(e,f). The final results of the inversion of the synthetic data wavefield in Figure 1a. Figure 3e contains the average of the slowness loci of iterations 4 and 5 (the dashed line) and the correct solution (the solid line). Except for differences in the velocity gradients near 2 and 5 km depth, the obtained and desired profiles are very similar. Figure 3f shows the slowness trajectory obtained by averaging the results of iterations 4 and 5 superimposed on the wavefield obtained by migrating the slant stack wavefield in Figure 1b with this function.

integral constraints (in the form of traveltimes) on the velocity-depth function. Amplitude information is an expression of the behavior of velocity gradients and so provides derivative constraints. In order to make use of such information the data must be recorded in such a way as to be able to recover true amplitudes.

The example above illustrates one approach to inversion, that of finding a model which is in some sense a single best fit curve to the data. The uncertainty in such a model is indicated in the present scheme by the width of the p-z image, and is directly derived from the frequency content and the time resolution in the original data.

Another approach to inversion concentrates on determining the envelope of all possible models that are consistent with a given data set (cf. McMechan and Wiggins, 1972; Bessonova et al., 1974, 1976). In wave-field inversion an explicit envelope can be determined by using a property of the migration algorithm. If a migration is performed with a velocity estimate that is too high at every depth, it will indicate depths that are everywhere too great. Conversely, if the migration velocity is consistently too low, consistently shallow depths are produced. We have also seen that a consistent bias produces convergence to two pseudo-stable states. A property of these states is that they are mutually inverse (i.e. migrating with either of the corresponding velocity functions produces the other). This idea has a number of implications. First, such a pair of mutually inverse functions constitute an envelope since one is everywhere too deep and the other is everywhere too shallow. Second, it implies that one need not expend effort on determining sophisticated criteria for locating the slowness locus when envelopes are the primary desired result (all that is required is any picking criterion that is consistently biased). Third, a good estimate of the desired $v(z)$ function can be obtained simply by averaging those of the two states. (An example of this was already presented above.) Finally, the potential now exists for complete inversion by a computer program since a simple criterion is sufficient and produces both an envelope and a reasonably accurate estimate of the velocity profile.

The interactive mode is likely to be the most viable form for practical application since it can avoid, through the introduction of human intuition

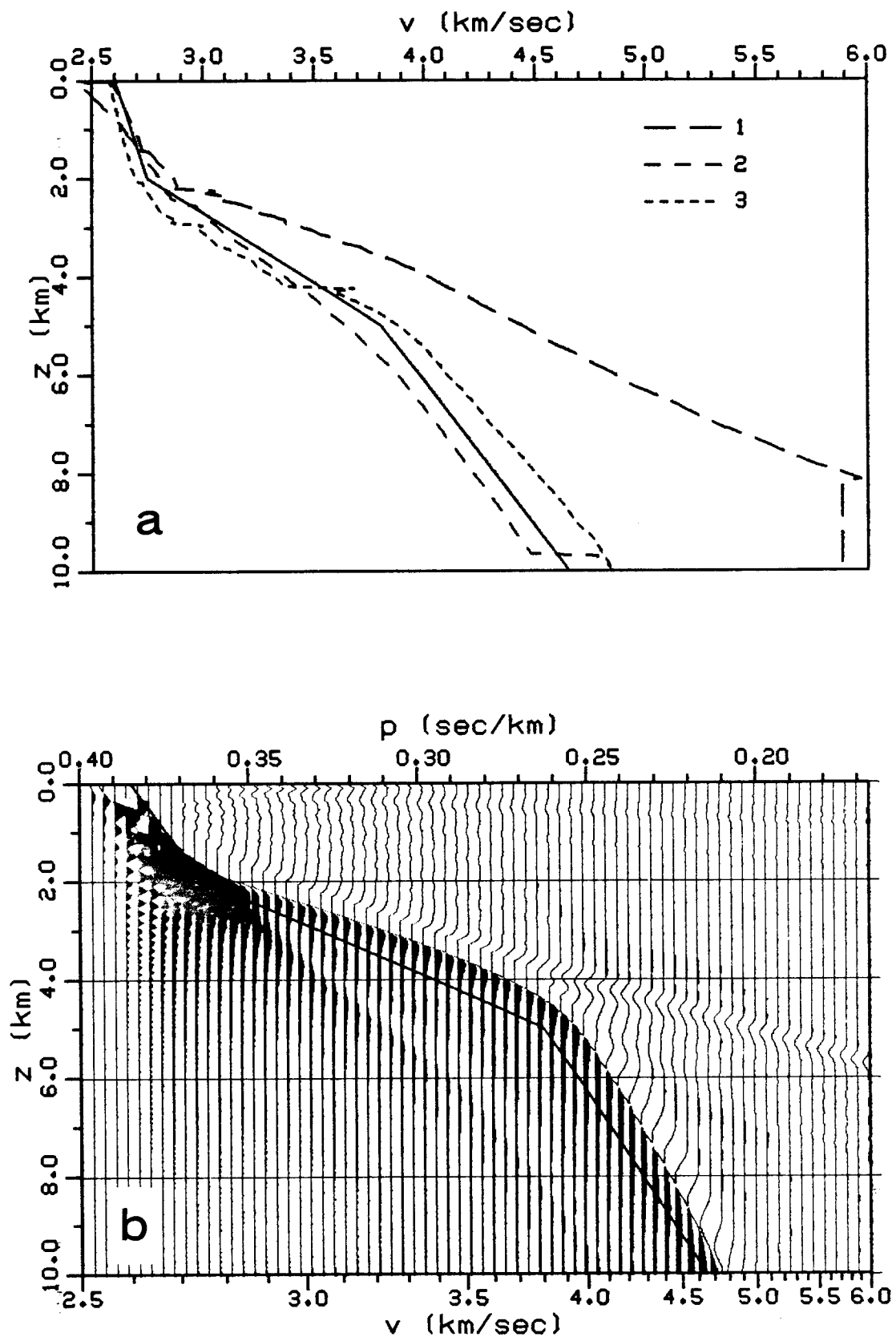


FIG. 4. Convergence in an automatic mode. The solid line in 4a is the velocity model for this example. The dashed lines are the results of the first three iterations, where at each step the slowness curve was picked automatically. The lower panel, 4b, shows the results of migrating with a velocity function (the solid line) that is the average of iterations 2 and 3.

and experience, most of the potential problems in picking the slowness trajectory from the migrated wave field. However, for certain datasets, it will be possible to do an inversion in a completely automated mode. Figure 4 contains an example of the result of an automated inversion. The solid velocity-depth profile in figure 4a was used to generate a synthetic refraction profile which was then slant-stacked. This p - r wave field was then iteratively migrated using a constant velocity of 2.6 km/s as the starting model. The picking criterion for finding a slowness locus was the maximum positive amplitude at each p . The velocity-depth functions produced by the first three iterations are shown in figure 4a as the dashed lines. Both iterations 2 and 3 are everywhere within 0.2 km/s of the correct solution. As a final estimate, the results of iterations 2 and 3 were averaged and a migration was done with this slowness function. The migrated wave field of this last iteration is shown in figure 4b. The solid line superimposed in figure 4b is the slowness function used for this iteration.

Figure 4b also illustrates the effect of the failure of the algorithm used to find the slowness trajectory that would compensate for the varying phase shifts on different arrival branches. Through the migration algorithm, an uncompensated phase shift at some z (which in effect produces a shifted slowness estimate) leads to an offset in z for all greater values of z . An error in picking at shallow z will be present at all greater z , and subsequent errors are cumulative.

Certain types of data, such as typical marine profiles and the Imperial Valley data presented below, contain prominent multiple reflections. Since these arrivals often mask primary arrivals their removal is the topic of much research. The transformations discussed here may be applied to this problem. Figure 5a contains the plane wave decomposition of a synthetic data profile consisting of both primary arrival branches and their first free-surface multiple branches. In the p - r plane, the energy in the multiple is separated from the primary energy since, for each p , it plots at twice the r of the primary. Similarly, in the migrated wave field in figure 5b, the primary and multiple images are separated. Thus, multiple removal may theoretically be performed by windowing the slant stack or migrated wave field to remove the multiple energy and then transforming back into the travelttime-offset domain to produce a multiple-free profile.

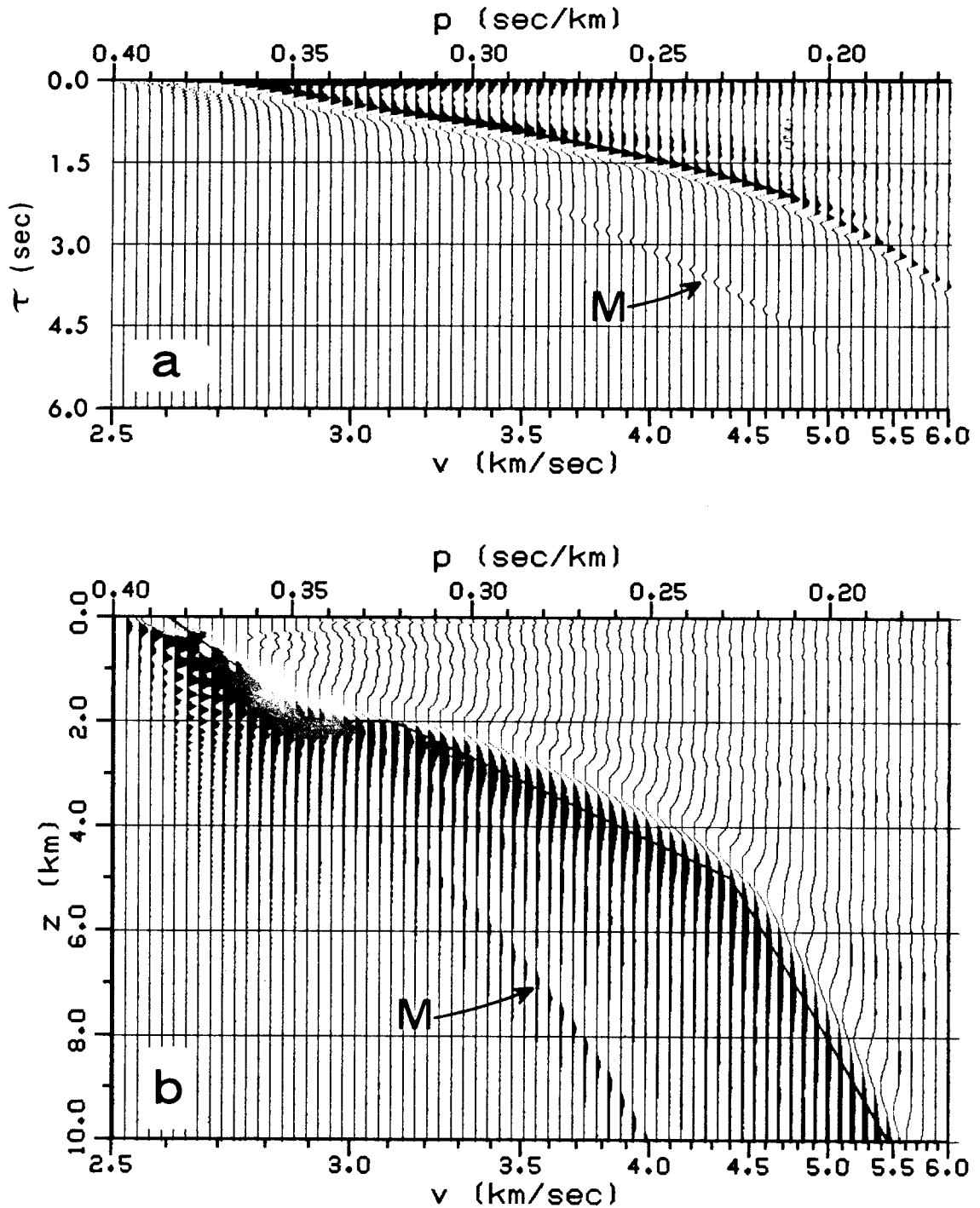


FIG. 5. The transformation of multiple reflections. First order multiples were included in the synthetic data profile used to produce the slant stack wavefield in (a). Migration of this wavefield with the correct velocity profile (the solid line in (b)) produces the wavefield in (b). In both (a) and (b) the image labeled M is the first multiple.

It should be noted that this approach to multiple removal is also valid for data collected over structures other than flat layers. The transformations themselves do not contain any restricting assumptions about the origin of the data (slant stacking does not involve any velocity assumptions, and even migration at a grossly incorrect velocity would still separate primaries and multiples). There are, however, still problems where a precritical reflection branch of a multiple crosses a primary branch. An alternate approach to multiples is to attempt to utilize the information present in them. We leave migration of multiples for future consideration.

One additional advantage that wave-field inversion possesses is that low velocity zones require no special treatment. The travelttime shadow zone associated with a low velocity zone (in the t - h domain) becomes transformed into a jump in τ at constant p in the p - τ domain and a gap in z at constant p in the p - z domain.

Finally, it is of theoretical interest to note that we have empirically shown that there is a direct correspondence between Wiechert-Herglotz inversion and migration. We expect that an asymptotic or stationary phase form of the migration equation (to delete the ω -dependence) can formally be shown to be equivalent to the Wiechert-Herglotz integral. Also, these concepts are expected to be applicable to the inversion of surface waves as well as body waves since it is possible, with certain assumptions, to invert a dispersion curve with the Wiechert-Herglotz integral (Takahashi, 1955; Nolet and Kennett, 1978). These topics are the subjects of current investigation.

Application to Recorded Data

In order to evaluate the applicability of wave-field inversion to actual field data, a digitized refraction profile of high spatial density was obtained from the United States Geological Survey (USGS). These data, which are presented in figure 6a, were recorded in the Imperial Valley of southern California in 1979.

The analysis of the Imperial Valley data followed closely the interactive procedure outlined above in the analysis of the synthetic data of figure 1a. The data were slant-stacked to produce the p - τ wave field shown in figure 6b.

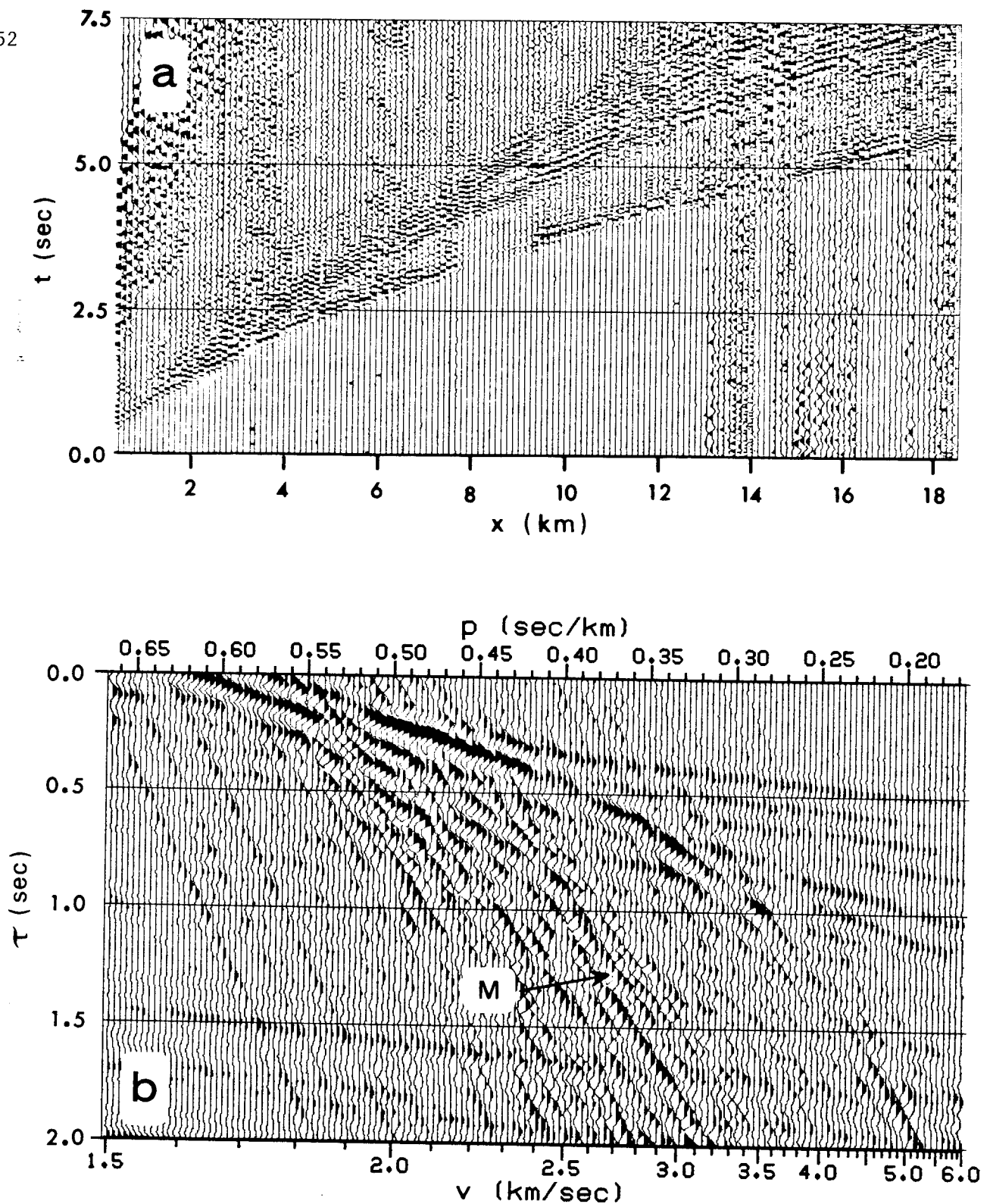


FIG. 6(a,b). Processing of real data. The record section in (a) was recorded in the Imperial Valley by the USGS. For this figure, each seismogram was scaled to have the same maximum amplitude and is plotted at approximately its correct offset (the actual spacing was not constant and there were a few bad traces that were not included). For the slant stack, the actual offset of each trace was used. The slant stack wavefield is shown in (b). The result of inversion of these data by migration is shown in Figures 6c and 6d. The image labeled M in (b) corresponds to the first multiple (PP).

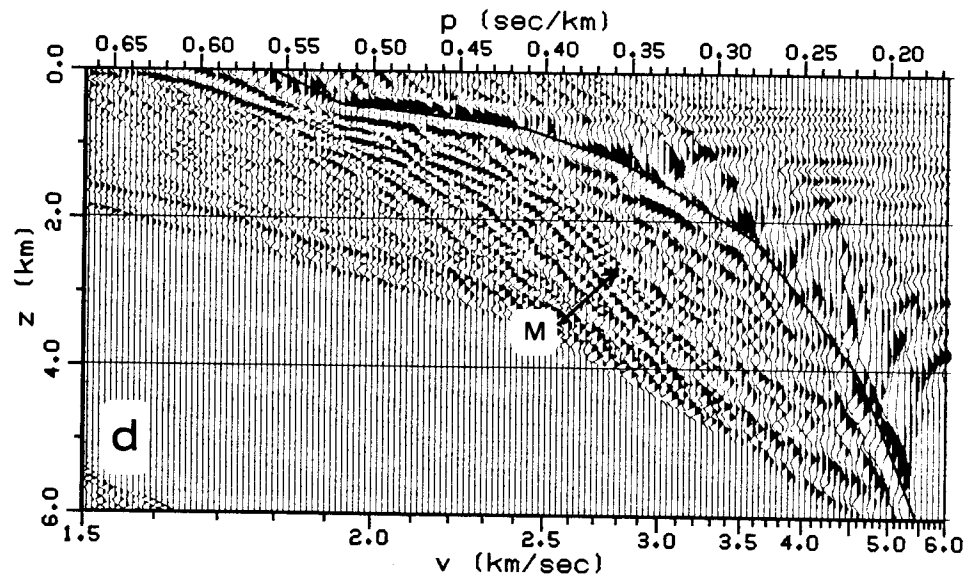
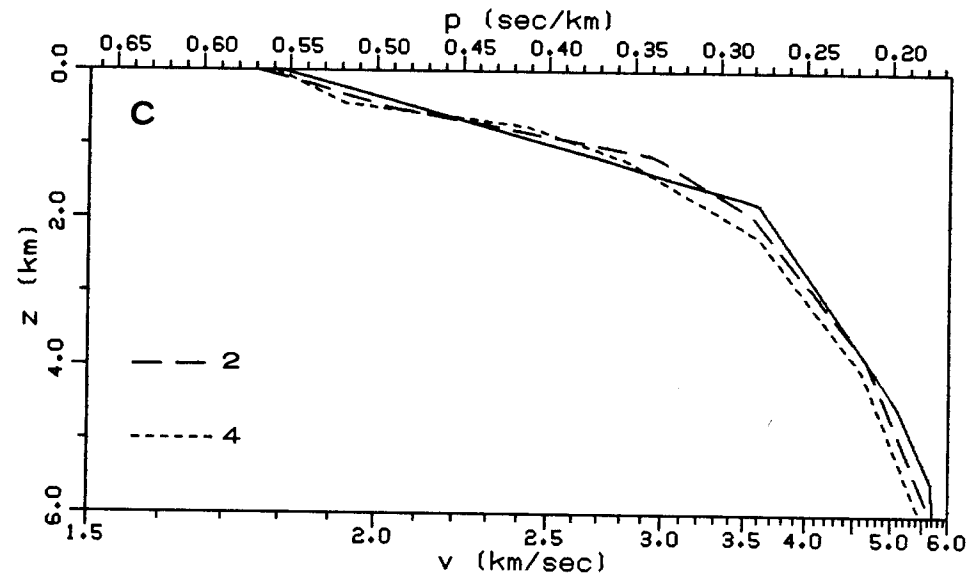


FIG. 6(c,d). The result of inversion by wavefield transformation of the data in Figure 6a. The two dashed lines in (c) are the slowness trajectories for iterations 2 and 4. The solid line is the result of inversion (by detailed ray tracing) of a reversed refraction profile near the line along which the data in Figure 6a were recorded. The lower (d) part of this figure illustrates the relationship between the slowness locus (the solid line) and the migrated wavefield at convergence. The image labeled M in (d) corresponds to the first multiple (PP).

At each iteration, the slowness trajectory was extracted by hand from the migrated wave field. Convergence was obtained in four iterations. In figure 6c, the velocity-depth profiles corresponding to the second and the final (the fourth) iterations, are shown. Also in figure 6c is the velocity profile obtained independently by W. Mooney of the USGS from a detailed ray tracing analysis of an earlier reversed refraction line in the same area. This previous line was approximately parallel to the one analysed here; the shot point for the later profile also corresponded nearly exactly to the previous shot point. Thus, it is reasonable to make a general comparison of the results of the two studies as presented in figure 6c.

Figure 6d contains the slowness trajectory extracted from the migrated wave field of the fourth iteration. This line is superimposed upon the wave field obtained by migrating with it to illustrate the convergence condition. The slowness locus is particularly clear in this wave field.

There are some features of this Imperial Valley example which illustrate points made in the theoretical discussion above. For example, the data (figure 6a) contain two regions of decreased resolution in the first arrival branch, one between 7 and 9 km offset where there is a decrease in amplitude, and one between 13 and 15 km offset where there is an apparent increase in the noise level. These two regions can be seen in transformed form in both the slant-stacked and migrated wave fields where they indicate an increased local uncertainty in r and z respectively. From the width of the slowness image, the depth resolution is estimated to be of the order of 0.1 km on the average, with better resolution at those depths associated with clear arrivals in the data and worse at those depths associated with noisy and less coherent arrivals. The prominent free-surface multiple PP, which can be seen in the data, is also transformed into the p - r and p - z domains as predicted in figure 5.

During the inversion of the recorded data the focussing observed with the synthetic data was very noticeable. As the correct velocity profile was approached, the sharpness of focus of the slowness trajectory approached that seen in the slant-stacked wave field.

In comparing figures 6b and 6d, there are some p - r loci which seem to migrate to horizontal lines. These are located near 1.5 km depth and may be

precritical reflections from small velocity contrasts, but we see no evidence for any large velocity discontinuities in this profile.

In this section the inversion of refraction data by wave-field transformations has been demonstrated by the processing of a real dataset from the Imperial Valley, California. These preliminary results encourage further application and development of the method.

Conclusions

In this paper we have presented an alternate method for inversion of spatially dense refraction data that is based on the technique of wave equation migration. The main advantage of the method is that the entire wave field is present throughout the inversion and the desired feature (the velocity-depth curve) is extracted directly from its image in the migrated wave field. This eliminates the subjective bias which can occur when traveltimes or tau curves are picked. The method is robust and is self-consistent in the sense that a consistent bias in extracting the image will produce convergence to a bistable state which envelopes the optimal solution. The depth resolution is indicated by the width of the p-z image and is directly derived from the frequency content and the time resolution in the original data. The method has been used, apparently successfully, to invert a dataset recorded in the Imperial Valley, California.

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